Kyoto Workshop, 23 February 2012

How much better do we understand the dynamics of stratospheric warmings – and what has it taught us about fundamental issues in geophysical and planetary fluid dynamics?

> Michael E. McIntyre, Dept of Applied Mathematics & Theoretical Physics, University of Cambridge, UK

Each question has the same quick answer, "lots". Talk will try to make both more specific:

- 1. Re warmings as such: recent advances & challenges. Whence linear theory?
- 2a. Fundamental points with wider implications, *e.g.* countering some myths about jets on Earth, on Jupiter, and in tokamaks. The overarching general-circulation paradigm change (history in Section 1 of Dritschel & McI (2008, *J. Atmos. Sci.).* Role of unconscious assumptions.

2b. Fundamental points that *should* be in the textbooks, *e.g.* Kelvin's circulation theorem ↔ "nonacceleration theorems" (shining exception: Bühler 2009)

Reprints, preprints & corrigenda: websearch "lucidity principles" then back to my home page at the strings "jets", "DIMBO", and "Rosenbluth"

 $2\pi\Omega/nHz$ convection 450 zone Schou et al '98 400 radiative interior – no significant evidence for departures from solid rotation (Princeton, May 2011) 350 tachocline Reprints, preprints & corrigenda: websearch "lucidity principles"

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Historical snippet: Here I quote myself against myself. In the 1982 review I wrote that Rossby wave breaking and the consequent PV mixing had perhaps escaped notice because of

## "the near-impossibility of drawing isentropic maps of potential vorticity" from data "and thus seeing directly what is going on".

I think I'd taken too much notice of the words of an eminent and intimidating colleague – contrary to our Royal Society's motto *Nullius in verba*.

Tim Palmer, Alan O'Neill and co-workers at the UK Met Office quickly proved me, and the eminent colleague, wrong with a **"damn fool experiment"**, computing mid-stratospheric isentropic maps of PV from satellite data giving us



"a blurred view of reality seen through... knobbly glass" of the "world's largest breaking waves":

Corrected (1984, J. Atm. Terr. Phys.) Clough et al 1985 (Q. J.. Roy. Met. Soc.)





"the near-impossibility of drawing isentropic maps of potential vorticity..."

Damn fool experiments at the UK Met Office























This was discovered accidentally in another "damn fool experiment"...

Edmon et al, J. Atmos. Sci. 37, 2600 and 38, 1115, especially 2nd-last item:



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The "saturation, propagation, saturation" pattern was a big surprise!

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The "saturation, propagation, saturation" pattern was a big surprise

- and it vindicated Dickinson's seminal 1969 work (JAS), the first dawning of a complete, robust understanding of Starr's "negative viscosity". (History in Dritschel & Mcl. 2008).

Hindsight shows the foregoing as episodes within a major paradigm change:



Understanding began to emerge **after** V.P. Starr's 1968 "negative viscosity" book, beginning with ←Jule Charney, Ernst Kleinschmidt, Bob Dickinson.→ History long and tortuous – no "Einstein moment". But **today's understanding is crystal-clear.** 



Historical review in Dritschel & McI (2008, J. Atmos. Sci.) on my home page.

The paradigm change (in our thinking about large-scale atmospheric dynamics, over the past century) can be summarized thus:

"turbulent atmosphere" (frictional)

"radiation-stress-dominated atmosphere" (often anti-frictional)

Accompanying insight: "there is no such thing as turbulence without waves."

Indeed, it's now clear that the generic role of wave propagation mechanisms illustrates one of the grand themes of physics, the **dynamical organization of fluctuations** with systematic mean effects.

But whence this phrase "damn-fool experiment"?

- Littlewood's Miscellany, ed. Béla Bollobás



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Some people seemed to think so at the time.

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... today the stratospheric "surf zone" is a hard-edged reality, familiar from advanced remote sensing and high-tech weather forecasting:

# McIntyre and Palmer (1983), revisited

#### Breaking planetary waves in the stratosphere

M. E. McIntyre<sup>\*</sup> & T. N. Palmer<sup>†</sup>

\* Department of Applied Mathematics and Theoretical Physics, University of Cambridge, Cambridge CB3 9EW, UK † Meteorological Office, Bracknell, Berks RG12 2SZ, UK

# 

## Movie



Fig. 2 Coarse-grain estimates of Ertel's potential vorticity Q on the 850 K isentropic surface (near the 10-mbar isobaric surface) on 17 (a) and 27 (b) January 1979, at 00 h GMT. The southernmost latitude circle shown is 20° N; the others are 30° N and 60° N. Map projection is polar stereographic. For units see equation (5) onwards. Contour interval is 2 units. Values greater than 4 units are lightly shaded, and greater than 6 units heavily shaded.

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## Final state



Fig. 2 Coarse-grain estimates of Ertel's potential vorticity O on the 850 K isentropic surface (near the 10-mbar isobaric surface) on 17 (a) and 27 (b) January 1979, at 00 h GMT. The southernmost latitude circle shown is 20° N; the others are 30° N and 60° N. Map projection is polar stereographic. For units see equation (5) onwards. Contour interval is 2 units. Values greater than 4 units are lightly shaded, and greater than 6 units heavily shaded.

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Our first view was indeed **blurred** and **knobbly**.

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Potential vorticity at 850K 00UTC 1979/01/27



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blurred and knobbly. Breaking planetary waves in the stratosphere Eddy-transport M. E. McIntyre<sup>\*</sup> & T. N. Palmer<sup>†</sup> barrier \* Department of Applied Mathematics and Theoretical Physics, University of Cambridge, Cambridge CB3 9EW, UK † Meteorological Office, Bracknell, Berks RG12 2SZ, UK Potential vorticity at 850K 00UTC 1979/01/27 а

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So the apparent success of linear theory is a **profound conundrum!!** 

Why / when does the linear theory of planetary-scale Rossby waves do so much better than it ought? (*Many* papers following Matsuno 1970, *J. Atmos. Sci.*)

This linear theory is **heavily relied on** in current thinking about the **annular modes** of variability – of **co-variability** of the stratosphere and troposphere – NAM and SAM –

– in particular, current thinking relies on the kind of linear behaviour studied in Chen and Robinson (1992, *J. Atmos. Sci.*) focusing on variability of the **Matsuno refractive index** near the subpolar tropopause.

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A quick reminder of those fundamentals:
"Suitably normalized" includes dividing by the mass of the fluid between adjacent stratification surfaces:



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PV invertibility is seldom flagged up clearly in textbooks - even when explaining Rossby waves!!

Rossby wave mechanism

#### (Q is the PV) Q high (cyclonic) Q anomalies (with sense shown by signs and curved arrows) Q anomalies (with sense (with sense (with sense (with sense (with sense (with sense (with sense) (with sense)

(~ material contours)

Q low (anticyclonic)







Only ideas needed are: PV invertibility & corollaries (including the Rossby-wave mechanism), and PV mixing by Rossby-wave breaking.



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Further insight comes from the **Taylor-Bretherton identity.** Shows how the whole jigsaw fits together, with PV mixing governing eddy momentum transport:

# The Taylor(-Bretherton) identity

(G.I. Taylor 1915, *Phil.Trans.Roy.Soc*; F. P. Bretherton 1966, *QJRMS*) It nonlinearly relates eddy fluxes of PV to momentum-flux divergences:

Barotropic (any 
$$L_{\rm D}$$
):  $\overline{v'q'} = -\frac{\partial}{\partial y} \overline{u'v'}$  (+ form stress if topog.)  
3D baroclinic:  $\overline{v'q'} = \frac{1}{\rho_0} \left( \frac{\partial F}{\partial y} + \frac{\partial G}{\partial z} \right)$   
where  
 $(F, G) = \rho_0(z) \left( -\overline{u'v'}, \frac{f_0 \overline{v'\theta'}}{N^2} \right)$  ("Eliassen-Palm flux")

**NB: nonlinear relation: valid at any amplitude!** And valid regardless of whether motion is free, forced, or self-excited. Often not flagged up clearly in textbooks. **My 1982 review just as guilty** – see p. 48b.<sub>5</sub>

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PV animation from recent numerical experiments by Gavin Esler illustrating the jet self-sharpening that's so typical and ubiquitous: Rossby waves

- (a) undulate the jet core elastically, and
- (b) break on both sides, mixing PV and sharpening the
  - jet's velocity profile (consequence of PV inversion)

The core acts as a remarkably effective "eddy-transport barrier" against mixing.

Esler, G., 2008, J. Fluid Mech. 599, 241



#### Strong jets, when disturbed naturally, tend to sharpen themselves. **A very simple toy model is insightful here:**

In this simplest model, the dispersion relation

 $c = u_{j} \left\{ 1 - \left( 1 + L_{D}^{2} k^{2} \right)^{-1/2} \right\}$ 

implies that the phase speed c lies within the range of jet velocity profile.

So rhe kinematics strongly favours Rossbywave breaking on the jet flanks. (This is the key message from nonlinear Rossby-wave critical-layer theory (Stewartson-Warn-Warn and beyond),

(The stratospheric examples are similar except that the polar-night jet self-sharpens mainly by PV mixing on its equatorward flank, as pointed out in my review.)



So why, then, **does the linear theory** of planetary-scale Rossby waves **do so much better than it ought?** 

Two likely (and inter-related) reasons:

1) PV inversion is insensitive to small scales. In particular, planetary-scale Rossby waves notice mainly the largest scales.

(2) the nonlinear effects are largely captured by the PV-mixing ansatz even though the mixing will usually be imperfect.

So observed PV gradients **implicitly incorporate some of the nonlinearity** through the weakness of surf-zone PV gradients; recall Matsuno 1970:



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Another question: progress in extending idealized perfect-surf-zone models like that of Esler and Scott (2005, J. Atmos. Sci.) to more realistic vertical structures?

(Contour-dynamics model with perfectlymixed surf zone, implying  $K^2 = 0$  there, hence perfect sideways reflection.)

With a realistic lower boundary condition the system has an "external" or "barotropic" mode. Even within linear theory this has its own built-in "upper reflector".

Model has strong **self-tuning resonance by vortex erosion**, with modest forcing (quasi-topographic) – big EP uprush!



The authors judge that the **wave-2 warming of February 1979** illustrates barotropic self-tuning resonance, but not the **SH wave-2 warming of September 2002**, which had more phase tilt leading to a double helix [*sic*]. And **SH final warming** (usually wave 1)? History reminds us how science is a struggle with unconscious assumptions. Here's another reminder (that we all make 'em):

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energetics assumption

small-is-unimportant assumption (counterexample: amplifier input signal)

eddy-viscosity assumption (counterexamples: Earth, Sun, laboratory)

homogeneous-turbulence assumption (and inverse-cascade assumption)

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One example is where A = B is the equation for zonal-mean temperature or potential temperature in a model of the Brewer-Dobson circulation: B = mean heating rate, A = (TEM residual) mean vertical velocity times the static stability of the stable stratification. The old idea that the right-hand side *B* can be regarded as **prescribed**, and the circulations as thermally **driven** is, indeed, just an A = B assumption.

But hang on – what's wrong with that? Why shouldn't I do a thought-experiment in which one prescribes the heating rate?

Answer: it's an **unnatural thought-experiment** in this context **because** the stratosphere and mesosphere (and the interiors of solar-type stars) are all **thermally-relaxing systems**. So it's more insightful to regard the heating rate *B* as part of the **response** to some forcing.

If I push a dinner-plate along a tabletop, then the friction force is part of the **response** to the force I apply. If I keep on pushing the plate, then it keeps on moving. If I stop pushing the plate, then it grinds to a halt. Its motion relaxes, frictionally, toward zero. (Would a thought-experiment **prescribing the friction force** make any sense?)

Similarly, with the stratospheric and mesospheric circulations, it's more insightful to say that they are driven not by heating but by wave-induced **(non-frictional) zonal forces** – mostly from breaking Rossby waves and breaking gravity waves. Keep on sending in the waves, and the circulation keeps going. It also tends to **burrow** – to extend itself **downward** from the forcing level (Haynes + 1991 *JAS*)

We may usefully describe these circulations as gyroscopically pumped.

"Einstein's Tealeaves" demonstrates gyroscopic pumping for the special case in which the zonal force happens to be frictional:



#### This experimental demonstration is very robust. It always works.

Take a cylindrical container with a rotating mass of fluid in it. The fluid near the bottom feels a retrograde frictional force. This fluid is **gyroscopically pumped** toward the centre. The tea-leaves follow it, as in Einstein's original example of flow in a teacup.

Jet mythology, zoology, physiology, and anatomy...

# Zoology:

- 1. Classic tropopause/polar-night/major-oceanic (Gulf-stream-like)
- 2. Mid-oceanic "striations" or "ghost jets", e.g. Maximenko et al (2008 GRL)
  - 3. Jovian jets (straight!)
  - 4. Tokamak jets (Marshall Rosenbluth Lecture, available on my home page)

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**Physiology** (starting point only):

Assume free or forced-dissipative **balanced vortex dynamics**  $\leftrightarrow$ PV invertibility principle holds. (So, e.g. **any** inhomogeneity that makes a background PV gradient even slightly step-like, or staircase-like, must give rise to jets. So we expect jets to be generic, whatever the mechanism(s).

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Anatomy (2 clear extremes):

Strong jets (PV-staircase-like, Rossby waves **guided**)

Weak jets (PV close to large-scale background beta, Rossby waves **unguided**, quasi-plane)

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Assume free or forced-dissipative **balanced vortex dynamics**  $\leftrightarrow$ PV invertibility principle holds. (So, e.g. **any** inhomogeneity that makes a background PV gradient even slightly step-like, or staircase-like, must give rise to jets. So we expect jets to be generic, whatever the mechanism(s).


The literature on jets – a complex conceptual landscape.

## Zoology:

- 1. Classic tropopause/polar-night/major-oceanic (Gulf-stream-like)
- 2. Mid-oceanic "striations" or "ghost jets", e.g. Maximenko et al (2008 GRL)
  - 3. Jovian jets (straight!)
  - 4. Tokamak jets (Marshall Rosenbluth Lecture, available on my home page)

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Not so clear: hyper-strong, hyper-staircase-like? Jupiter? (Dowling 1993, JAS)

Phillips Effect



Phillips Effect







(PV inversion then gives jets.)



So if PV mixing occurs, it tends to be spatially **inhomogeneous**. (PV inversion then gives jets.) Feedback stronger in strong-jet cases: PV inversion implies **reinforcement by shear** to form a classical **eddy-transport barrier** (Juckes & M, *Nature* 1987).



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**PV map** and **dye map** near-identical.

This **is** clearly a **strong jet:** staircase-like; eddy-transport barrier.



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**PV map** and **dye map** near-identical.

This **is** clearly a **strong jet:** staircase-like; eddy-transport barrier.

By the way: no upscale cascade is involved. (Surprise??)



Model stratospheres are similar (Juckes & M 1987):

Polar-night jet strengthened and sharpened by PV mixing mainly on its equatorward flank, forming a **strong jet** and **eddy-transport barrier** 

(This is a well-studied problem!)

Again, no upscale cascade.



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Again, no upscale cascade.

Same for the **real** stratosphere.





Summary: 2-level hierarchy of ideas for understanding the fluid dynamics of jets

1. Generic ideas:

PV Phillips effect

Taylor-Bretherton identity  $\overline{v'q'} = - \operatorname{div} (\operatorname{eddy} \operatorname{momentum} \operatorname{flux})$ Nonlinear, forced/free/self-excited

### 2. Particular mechanisms:

i) Rhines effect. Re weak jets generated by strong small-scale forcing – strong enough to create active small-scale vortices that merge or cluster, producing an upscale cascade that is arrested or slowed when eddy velocities ~ plane Rossbywave phase speeds. Wave-turbulence interaction is spatially homogeneous.

(ii) Jet self-sharpening by Rossby-wave breaking. Re jets **strong** enough to be Rossby waveguides. Wave-turbulence interaction spatially

inhomogeneous. (iii) Repeated excitation of Kelvin sheared disturbances by small-scale forcing weaker than in (i). (Kelvin 1887, Farrell and Ioannou 2007 & refs.).

(iv) Downstream wind stress reinforcing strong ocean jets (e.g. Thomas & Lee'05 JPO)

Reprints, preprints & corrigenda: websearch "lucidity principles" then back to my home page at the strings "jets", "DIMBO", and "Rosenbluth"

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MYTH: "Mechanism (i) is universal."

Stratosphere disproves this: clearly (ii). For Jupiter I'm betting on (iii) .

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#### **DIMBO** = **DI**apycnal **M**ixing via **B**aroclinic **O**verturning



How deep is the DIMBO layer? Scale analysis and semigeostrophic PV inversion suggest the "obvious" answer fL/N. Could ~ kilometre or two. Must often exceed mixed-layer depth.

Numerical experiments underway (John Taylor, Raff Ferrari, personal communication) – watch this space!

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The old idea that the right-hand side *B* can be regarded as **prescribed**, and the circulations as thermally **driven** is, indeed, just an A = B assumption.

But hang on – what's wrong with that? What's wrong with a thought-experiment in which one prescribes the heating rate?

Answer: it's an unnatural thought-experiment in this context **because** the stratosphere and mesosphere – and the interiors of solar-type stars – are all **thermally-relaxing systems**. So it's more insightful to regard the heating rate B as part of the **response** to some forcing.

If I push a dinner-plate along a tabletop, then the friction force is part of the **response** to the force I apply. If I keep on pushing the plate, then it keeps on moving. If I stop pushing the plate, then it grinds to a halt. Its motion relaxes, frictionally, toward zero. (Would a thought-experiment **prescribing the friction force** make any sense?)

Similarly, with the stratospheric and mesospheric circulations, it's more insightful to say that they are driven not by heating but by wave-induced **(non-frictional) zonal forces** – mostly from breaking Rossby waves and breaking gravity waves. Keep on sending in the waves, and the circulation keeps going. It also tends to **burrow** – to extend itself **downward** from the forcing level (Haynes + 1991 *JAS*)

We may usefully describe these circulations as gyroscopically pumped.

# Gyroscopic pumping is **easy** to understand:

Rapidly-rotating system! Low Rossby number, Coriolis effects are **strong**.

Coriolis force turns fluid poleward: a robust and systematic **mechanical pumping effect**. Another example is **Ekman** pumping.



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#### control"



#### This experimental demonstration is very robust. It always works.

Take a cylindrical container with a rotating mass of fluid in it. The fluid near the bottom feels a retrograde frictional force. This fluid is gyroscopically pumped toward the centre. The tea-leaves follow it, as in Einstein's original example of flow in a teacup.

# But how does a thermally relaxing, stably-stratified system respond to gyroscopic pumping at some altitude? Ans: "downward control"

Terrestrial cases with a lower boundary (Haynes, P. H., et al., 1991):





Immediate consequence of mass conservation together with the "gyroscopic-pumping relation"

$$\frac{\partial \psi}{\partial \phi} \frac{\partial \bar{m}}{\partial z} - \frac{\partial \psi}{\partial z} \frac{\partial \bar{m}}{\partial \phi} = \rho_0 a^2 \bar{\mathcal{F}} \cos^2 \phi$$
(hyperbolic operator!)  
Haynes, P. H., et al., 1991: J. Atmos. Sci., **48**, 651-678

When force is due entirely to upward-propagating gravity waves, and Rossby number small, there is an interesting simplification, relevant to the cold summer mesopause: E.g., **polar mesospheric upwelling** depends only on gravity-wave flux from below:  $\bar{w}_{0}^{*}(u, z) = -\frac{1}{2} \frac{\partial}{\partial z} \{f^{-1}(u) \cos u \, \overline{u' u'}\}$ 

 $ar{w}_0^*(y,z) = - rac{1}{\cos arphi} rac{\partial}{\partial y} \{f^{-1}(y) \cos arphi \ \overline{u'w'}\}$ 

(McIntyre, JGR 1989, special issue on noctilucent clouds; also Shepherd and Shaw JAS 2004) June 75 °N



Tuning Stratosphere Some nostalgia depends from the early and m QBO? 1980s: troposphere together Response Tropospheric forming a AFFECTS anomalies on NONLINEAR hemispheric-TUNING time-scales of weeks (Charney-DeVore COUPLING scale cavity Plumb) at ultra-long ("blocking") reinforcing wavelengths waves 3, 4, 5, topog. Eslor Scot • Equivalent-barotropic \* 2005 JAS RESPONSE dynamics ? Storm-track forcing ? depends on · Linearize about tuning of cavity time-mean, Emally-asymmetric state? et assu "artificial (Simmons, Branstator, Wallace) 1 surf-zone · Local resonance ? (15 compatible. encroachment: (see my JMSJ review) with insta (1982)



125

3.0











T. DUNKERTON, C.-P. F. HSU AND M. E. MCINTYRE JAS 1981 827



**NB:** response **TALLER** than forcing by EP flux div'gce  $D_F$ 



Some questions from my 1982 JMSJ review:

1. How and why do planetary-wave amplitudes become anomalously large?

2. When they do become large, for what stratospheric conditions is a major warming likely to occur (and why are major warmings relatively uncommon)?

3. To what extent can we use linear planetary-wave theory for the wave structure? And in particular,

4. to what extent can we think of the principal zonal wavenumbers, 1 and 2, as acting independently of each other? In other words, how much can we explain without invoking nonlinear interactions between different zonal wavenumbers? 5. Are wave-reflection and resonance phenomena important or not (*e.g.* to question 1)?

6. Are "critical lines" important or not (e.g. to question 2, or to question 5)?

7. Are shear instabilities involved at any stage, and are they relevant to question 1?

8. What quantities should be monitored in order to be able to forecast warmings?

9. To what extent, and in what sense, does the troposphere behave independently of the strato-sphere (for the purposes of question 1 for instance), and how should we represent tropospheric-stratospheric coupling in mechanistic models?

Further questions include whether the Antarctic final warming can be regarded as involving wave-1 resonance...

Esler-Scott work now makes the clearest case for a resonant response mode **inherently involving the stratosphere and troposphere tightly coupled together**...

A **BIG** remaining problem is how to quantify the coupling, e.g. via tropo-Spheric eddy fluxes in storm tracks and other features tied to geography Early 1980s... a story of scientific good luck: the stratosphere as an outdoor fluid-dynamical laboratory...

Significant events visible because of their large vertical scale, visible even to nadir sounders...

Beautiful early case study: Clough, S. A., Grahame, N. S., O'Neill, A., 1985: Q. J. Roy. Meteorol. Soc.,111, 335-358
Courtesy Dr A J Simmons, European Centre for Medium Range Weather Forecasts:

# McIntyre and Palmer (1983), revisited

### Breaking planetary waves in the stratosphere

M. E. McIntyre<sup>\*</sup> & T. N. Palmer<sup>†</sup>

\* Department of Applied Mathematics and Theoretical Physics, University of Cambridge, Cambridge CB3 9EW, UK † Meteorological Office, Bracknell, Berks RG12 2SZ, UK

# 

### Initial state



Fig. 2 Coarse-grain estimates of Ertel's potential vorticity O on the 850 K isentropic surface (near the 10-mbar isobaric surface) on 17 (a) and 27 (b) January 1979, at 00 h GMT. The southernmost latitude circle shown is 20° N; the others are 30° N and 60° N. Map projection is polar stereographic. For units see equation (5) onwards. Contour interval is 2 units. Values greater than 4 units are lightly shaded, and greater than 6 units heavily shaded.

-100 0 100 200 300 400 500 600 700 800 900 1000 1100 1200 **PVU**  I had been too impressed by the words of an eminent colleague: "anyone who tries to compute so highly differentiated a quantity as PV from observational data **is a fool.**"

## from old savannah talk:

(so, e.g., PV mixing has to **fit in** with radiation stress: this is **why Rossby-wave breaking can mix PV** so easily)

(BUT Jupiter might be a case of stress governing PV flux...

Another question: any progress in extending idealized perfect-surf-zone models like that of Esler and Scott (2005, J. Atmos. Sci.) to more realistic vertical structures?

Contour-dynamics model with perfectly mixed (therefore perfectly reflecting) surf zone

With a realistic lower boundary condition the system has an "external" or "barotropic" mode. Even within linear theory this has its own built-in "upper reflector".

Can get strong **self-tuning resonance by vortex erosion**, with modest forcing (quasi-topographic) – big EP uprush!

The authors judge that the wave-2 warming of February 1979 involved simple barotropic resonance, but not the SH wave-2 warming of September 2002., which had more phase tilt leading to a 2-vortex helix.



Indeed, the dynamics is **strongly** nonlinear. There are intimate, promiscuous, non-resonant interactions among many Fourier components. It's also spatially **very inhomogeneous.** 

Homogeneous turbulence theory is **inapplicable (!)** 

And standard theoretical physics might well say that the problem of understanding the dynamics is hopelessly intractable. **However**,

some of the most important aspects are captured by a very simple, yet powerful idea -- the idea of **PV mixing** (complete or partial).

(It's an old idea, by the way. Its ancestry is traceable back to G. I. Taylor's vorticity-mixing ideas, 1915 onward).

Why is this idea powerful? Let's remind ourselves of the main properties of the PV.

It's useful to define the exact PV in a way that applies equally to single-layer and multi-layer, compressible and incompressible, stratified rotating systems:

Momentum transport is intimately associated with PV mixing. (Material invariance says "mixing" can make sense.) Two ways to see it:

# (2) Idealized thought-experiments on PV mixing.

These idealize phenomena that are observed for real: Aug. 10, 1997

CRISTA N<sub>2</sub>O in upper stratosphere, courtesy Martin Riese



This is the so-called "stratospheric \_ surf zone"

For a tutorial on the dynamics,

websearch "dynamics that is significant for chemistry"